

# Mesopause Region Temperatures and Plasma Scale Height Estimations from VHF Meteor Radar and LF Absolute Reflection Height Measurements at Collm

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## Abstract

The change of ionospheric absolute reflection heights  $h$  of low-frequency (LF) radio waves at oblique incidence in the course of the day is measured at Collm Observatory (51.3°N, 13.0°E) using 1.8 kHz sideband phase comparisons between the sky-wave and the ground wave of a commercial 177 kHz transmitter (Zehlendorf, reflection point 52.1°N, 13.2°E). Plasma scale height estimates  $H$  are calculated from the decrease/increase of  $h$  in the morning/evening. The day-to-day variations of  $H$  are compared with those of daily mean temperatures at 90 km, measured with a VHF meteor radar (36.2 MHz) at Collm utilising the amplitude decay of meteor reflections. A good qualitative correspondence is found between the two data sets. Since mesospheric long-period temperature variations are generally accepted to be the signature of atmospheric planetary waves, this shows that LF reflection height measurements can be used for monitoring the dynamics of the upper middle atmosphere.

## 1 Introduction

The dynamics of the mesosphere/lower thermosphere (MLT) region at 80-100 km altitude is forced through wave coupling, so MLT mean circulation and temperature trends indicate possible long-term trends in waves and vice versa. Usually, waves are distinguished between short-period gravity waves, tides with periods of 24 hours and harmonics, and other planetary waves in the so-called long-period range up to about 30 days.

Planetary wave analyses often describe either the total variability of winds or temperature in the period interval to about 30 days (Jacobi et al., 1998; Bittner et al., 2000), or they focus on one or few specific waves as, e.g., the Rossby normal modes as the quasi 2-day wave (Chshyolkova et al., 2005). Only few analyses of waves on a global or hemispheric scale are available, and none of them considering interannual variability. Regarding long-term trends and possible indicators for climate variability, there is a tendency for a slight increase of overall day-to-day variability in the 1980s and 1990s (Jacobi et al., 1998, Bittner et al., 2000), but this result is based on only few datasets, and is to date neither generally confirmed by a global analysis, nor explained through theoretical consideration. A more indirect analysis of planetary waves was performed using radio wave absorption variations (e.g., Lastovicka et al., 1994). These analyses showed a possible positive trend, but this is intermittent (Lastovicka, 2002).

To summarise, the current knowledge on planetary wave activity in the MLT is still incomplete, mainly because of the lack of long homogeneous time series. To contribute to the overall data base and providing a potential for additional wave analysis, we propose here to

use the diurnal change of lower E region low-frequency (LF) reflection heights as an indicator for daily temperatures and to derive potential waves from these analyses. Using LF radio waves as a tracer for synoptic measurements has been done since several decades (Entzian et al., 1976; von Cossart and Entzian, 1976; Lauter et al., 1977). Usually phase-height measurements utilise LF transmitters located several hundreds of kilometres away from the receivers, with reflection heights well below 85 km during daytime. Here, we make an attempt to analyse scale heights using LF waves on a short propagation path.

## 2 LF height measurements and analysis of scale height

Low frequency 177 kHz radio waves from a commercial radio transmitter are registered at Collm Observatory, Germany (51.3°N, 13.0°E, distance to transmitter 170 km). The virtual reflection heights  $h'$ , referring to the reflection point at 52.1°N, 13.2°E, are estimated using measured travel time differences between the ground wave and the reflected sky wave through phase comparisons on sporadic oscillation bursts of the amplitude modulated LF radio wave in a small modulation frequency range around 1.8 kHz (Kürschner et al., 1987). The reflection height measurements have started in September 1982 and are carried out continuously since then, even after the temporary change to the digital radio broadcasting system DRM in September 2005.

The height resolution of an individual reflection height measurement is nearly 2 km. The data are combined to half-hourly means that consist of approximately 6000 individual values on an average. Essentially caused by the variability of reflection heights the 1- $\sigma$  variation of the half-hourly mean is in the order of 3 km below 95 km and somewhat above 5 km near 100 km height.

Half-hourly virtual reflection heights  $h'$  range between roughly 82 km during daytime and nearly 100 km during nighttime for the ordinary component of the electromagnetic sky wave

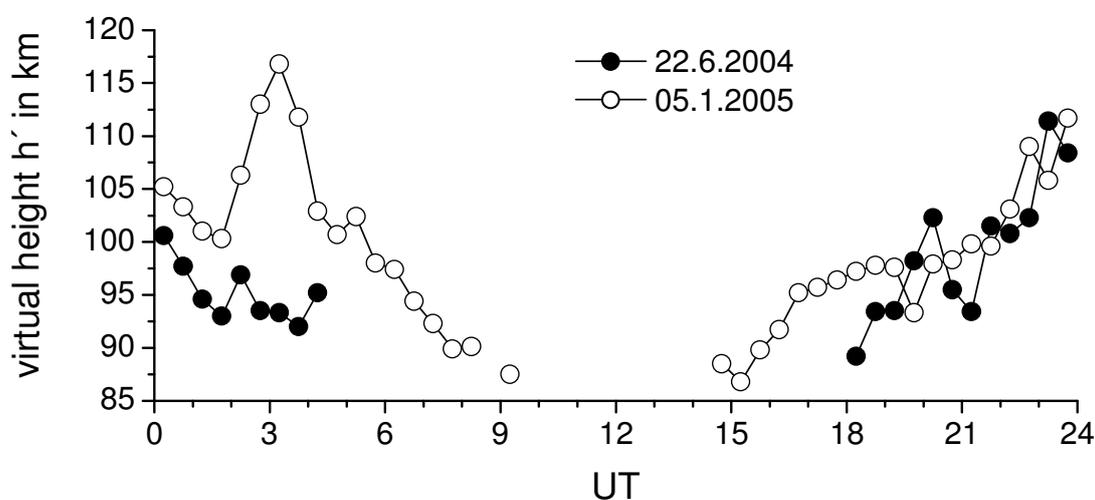


Figure 1. Two arbitrarily chosen examples of virtual height variation in the course of one day during summer and winter.

field. Especially in winter in the late night and early morning hours and in some midsummer nights the wave propagation is subject to stronger magnetoionic splitting into the ordinary and extraordinary component, with very large virtual reflection heights of the latter. Note that in this case the half-hourly mean height consists of a mixture of both components with unknown weighting, and usually will not be used here. During daylight hours, particularly in the summer months height measurements are not possible due to the strong D-region absorption of sky waves then. This is especially the case with the summer ionosphere. Figure 1 shows examples of available data on a summer and a winter day. It can be seen that in summer the daytime decrease of  $h'$  cannot be detected well, and its analysis requires additional effort. Therefore in this first step we restrict ourselves on the analysis of winter months.

Assuming essentially equilibrium between ionisation and recombination, a one gas and isothermal atmosphere, monochromatic radiation, and with that resulting a Chapman profile of the electron density, the height  $h$  of the layer maximum changes with zenith angle  $\chi$  as:

$$h = h_0 + H \ln \text{Ch}(\chi), \quad (1)$$

with  $h_0 = h(\chi = 0)$ ,  $H = RT/Mg$  as the plasma scale height ( $T$  as absolute temperature,  $R$  as molar gas constant,  $g$  as constant of gravitation and  $M$  as molecular mass) and the Chapman function  $\text{Ch}(\chi)$ . If  $\chi < 75^\circ$  then  $\text{Ch}(\chi) \approx 1/\cos\chi$  is a suitable approximation in the case that the curvature of the earth is neglected. Experimentally, a similar relation has been found between the virtual reflection height  $h'$  of LF waves below the layer maximum and the solar zenith distance as in the case of Chapman layer maximum:

$$h' = h'_0 + H \ln \text{Ch}(\chi). \quad (2)$$

Therefore, plotting the logarithm of the Chapman function against the virtual reflection height should provide as a first approximation a nearly linear correspondence, with the slope of the regression line as a scale height estimate:

$$H = \frac{\Delta h'}{\Delta \ln(\text{Ch}(\chi))}. \quad (3)$$

In the case of a delay of the reflection height curve with respect to the solar zenith distance, i.e. for the case of non-equilibrium, the slope differs from  $H$ . While von Cossart (1976) for used the average of forenoon and afternoon mesospheric measurements and obtained a good correspondence with rocket measurements, the delay in the lower E region may differ from those at lower heights. Moreover, the difference between real and virtual reflection heights can be substantial and reach more than 2 km in the D region and more than 5 km in the lower E region. Therefore, the results of  $H$  estimations have to be considered as qualitative. In addition, at higher altitudes the derived  $H$  values are smaller, especially in winter (Entzian et al., 1976, Lauter et al., 1977), which is known as the so-called "winter anomaly". Caused by the comparatively short distance between transmitter and receiver, which is connected to a steep ionospheric incidence angle the mean reflection height on 177 kHz lies above that level in earlier works (75 - 95 km) (Lauter et al., 1966). Thus, differences between forenoon and afternoon just as summer and winter effects on the  $H$  estimates are to be expected. As an example, in Figure 2 for two months  $h'$  vs.  $\ln(\text{Ch}(\chi))$  is shown separately for the forenoon  $h'$  decrease (open down triangles) and the afternoon  $h'$  increase (solid up triangles). Median

values for these two cases (squares for forenoon, circles for afternoon) are given in the figure, and slopes are calculated. Clearly, the slopes are different and the change of slope for larger  $\ln(\text{Ch}(\chi))$  values, i.e. for low elevation angles, can also be seen. Therefore, slopes have been calculated using the forenoon and afternoon median values only in the range of  $\ln(\text{Ch}(\chi)) < 3$ . For the afternoon the resulting slope is clearly too small, while the values for the forenoon are more realistic. In the following we shall use forenoon data only.

Figure 2 also shows that sometimes at small values of  $\ln(\text{Ch}(\chi))$  large values of  $h'$  are measured. These are owing to reflections from the extraordinary component mixed with the ordinary component. These data cannot be used for the analysis; in Figure 2 they are automatically discarded since we take the median of two months of data, but on a single day, as is exemplarily shown in Figure 3 these data can completely alter the profile. As shown in the figure, data that are obviously dominated by the extraordinary component are discarded. It should be noted that, for single days, therefore the analysis of the scale height is done from only very few data points and large uncertainty resulting from that.

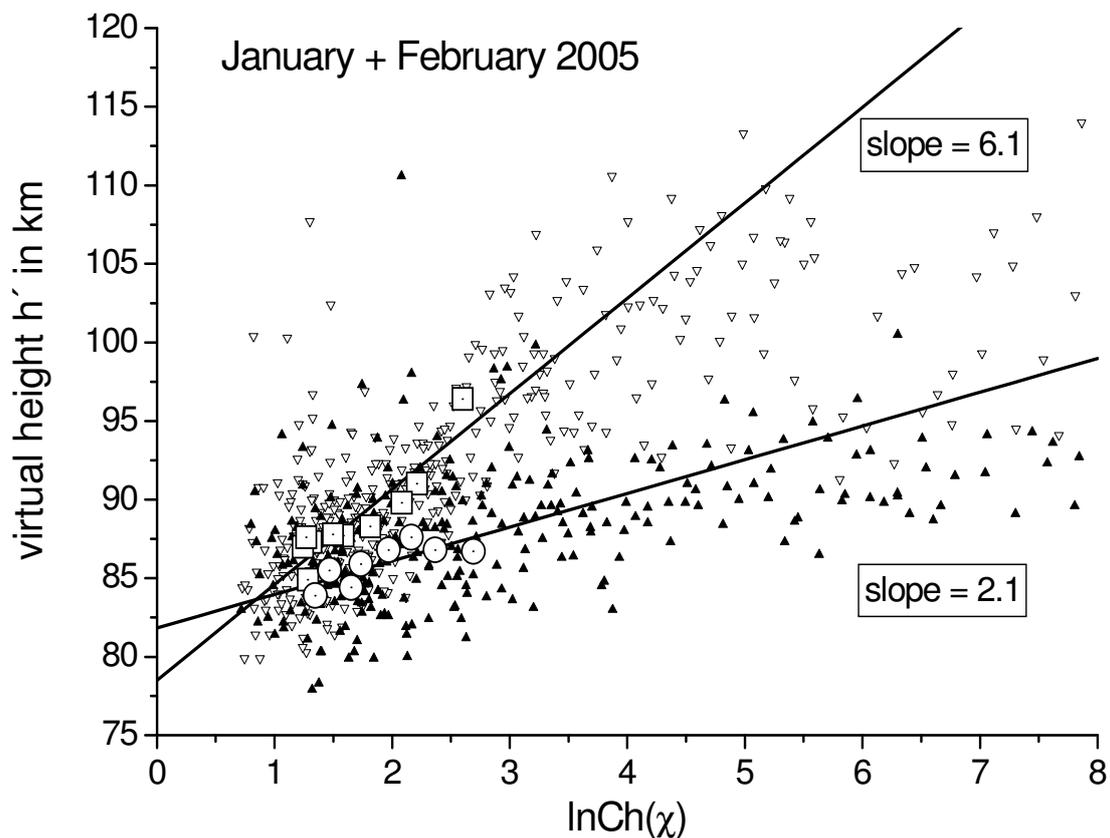


Figure 2. Virtual height  $h'$  measured on 177 kHz vs. the logarithm of the Chapman function  $\text{Ch}(\chi)$ , for each data in January and February 2005. Open down triangles show forenoon data, while solid up triangles show afternoon data. Median values of  $h'$  are given for  $\text{Ch}(\chi) < 3$ , and for forenoon (squares) and afternoon (circles) data separately. The slopes are calculated from the median values.

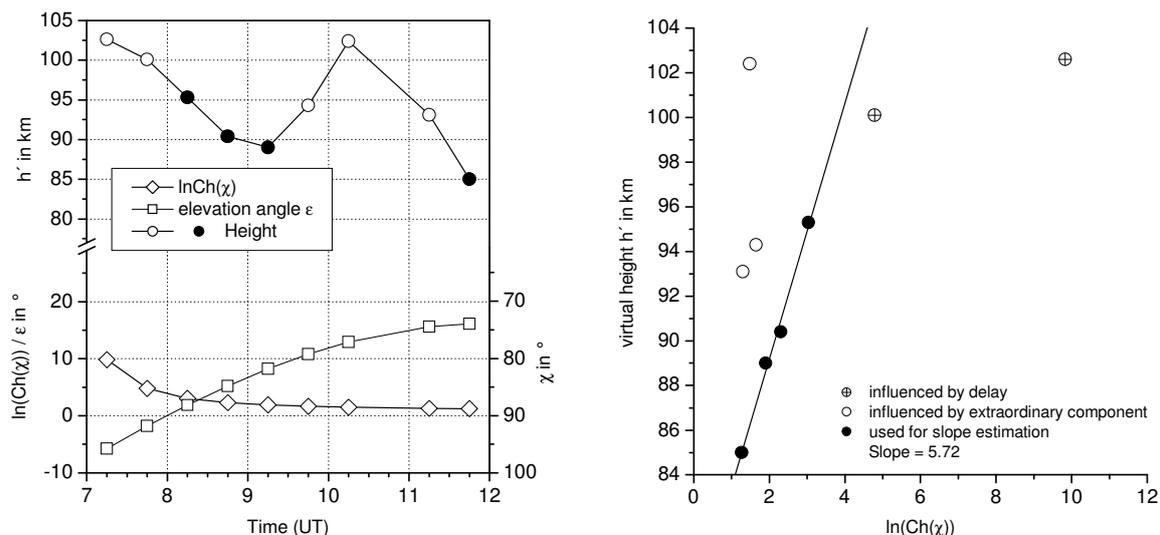


Figure 3. Example of scale height  $H$  estimation for the forenoon of 10 January 2005. Left panel: time series of logarithm of Chapman function  $\ln Ch(\chi)$ , elevation angle  $\varepsilon = 90^\circ - \chi$ , and virtual height  $h'$ . Right panel: virtual height  $h'$  vs.  $\ln Ch(\chi)$ . Solid dots denote  $h'$  data points used for  $H$  estimation.

### 3 Mesopause temperatures derived from VHF meteor radar measurements

VHF meteor radars measure the radio wave reflection from the ionised trails of meteors entering the Earth's atmosphere. The decay time of the signal is detected from so-called underdense meteor trails, i.e. from those trails whose reflectivity is determined by their electron density (Hocking, 1999, Hocking et al., 2001), which decreases with time due to diffusion. The amplitude decreases exponentially depending on the ambipolar diffusion coefficients  $D_a$ :

$$A = A_0 \exp\left\{-\frac{16\pi^2 D_a}{\lambda^2} t\right\}, \quad (4)$$

with  $A$  as the signal amplitude,  $\lambda$  as the radar wavelength and  $t$  as time. Measuring the decay of the amplitude allows to estimate  $D_a$ . The diffusion coefficient is proportional to the ratio of temperature squared and pressure. Introducing a constant temperature gradient from an empirical model and using exponential pressure decay with height, one obtains an equation for the temperature in dependence of the diffusion coefficient change with height:

$$\frac{d \ln D_a}{dz} = -2 \frac{dT/dz}{T_0} + \frac{mg}{kT_0}, \quad (5)$$

to be solved for  $T_0$ . Details can be found in Hocking (1999). Temperatures estimated with this method have been presented by Hocking et al. (2004) and Singer et al. (2004), also showing validation with rocket and ground-based optical methods. Temperature data are available as

daily means at the height 90 km, which is approximately the height where maximum meteor rates are found.

At Collm Observatory, a SKiYMET meteor radar is operated on 36.2 MHz since summer 2004. From meteor reflections mesopause region hourly wind profiles and daily temperatures are derived that roughly apply to a circle around Collm with 300 km diameter, and for heights between 80-100 km. First results using this radar and a more detailed description of the system have been provided by Jacobi et al. (2005).

#### 4 Results and discussion

The time series of meteor radar derived daily temperatures in early 2005 is shown in Figure 4. The daily scale height estimates after Eq. (3) are added. Both curves have been subjected to a 3 point FFT filter (thick solid line) to suppress fortuitous changes from the measurement and analysis uncertainty. The temperature and H curves show the same qualitative tendency. Apart from a short time interval around January 27, when, however, the H values are unrealistically large anyway, the correspondence is striking. Obviously, the dominating day-to-day and long-period changes are visible in both datasets, indicating that their variability is of the same origin, i.e. the MLT temperature change. Power spectra of both parameters are shown in Figure 5. Peaks are correspondingly found for periods of 3, 5, and 8 days. These are periods typical for planetary waves in the MLT.

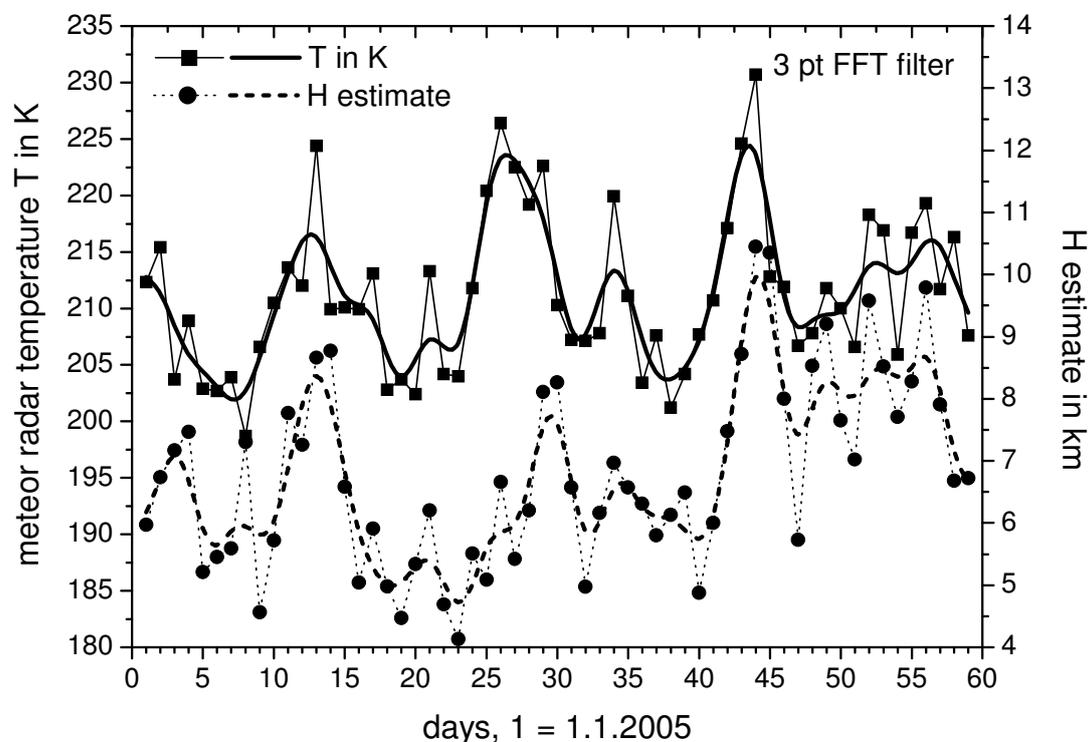


Figure 4. Time series of scale height  $H$  estimates and meteor radar temperatures  $T$  during January and February 2005. The thick solid lines are smoothed data using a 3 point FFT filter.

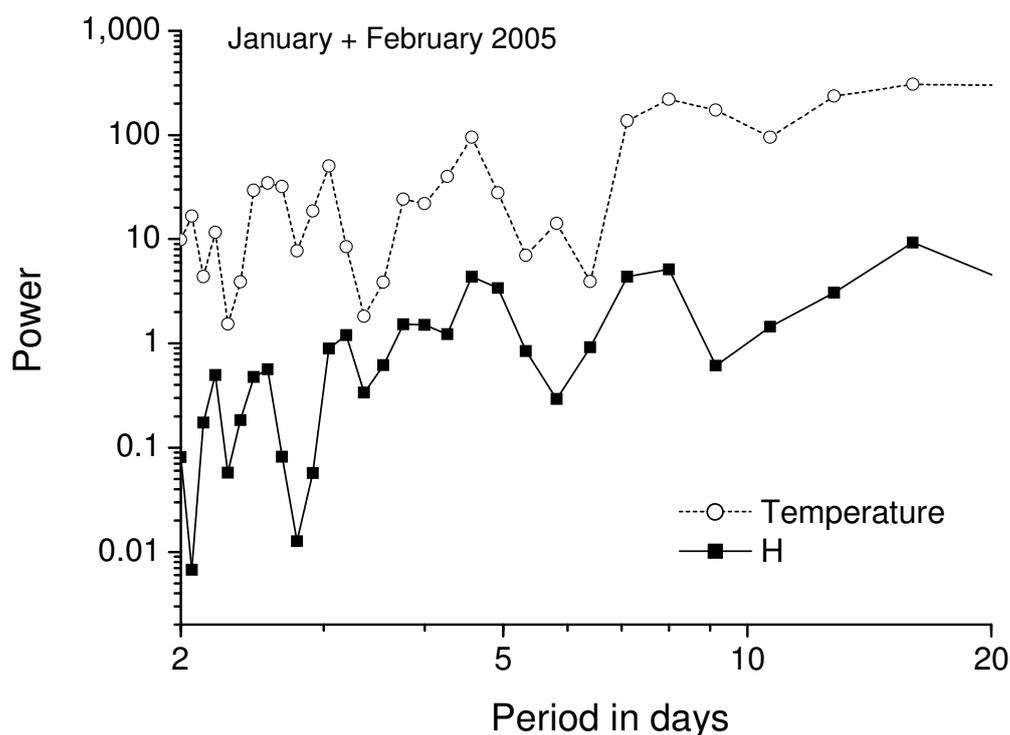


Figure 5. *Power spectra of scale height estimates and meteor radar temperatures.*

A comparison of scale heights derived from meteor radar measurements and LF plasma scale heights is shown in Figure 6. The figure clearly shows the qualitative nature of a single  $H_{LF}$  estimation. The scatter range of values is much larger than the scale height estimation from the radar. This is in qualitative agreement with results of von Cossart and Entzian (1976) for the upper mesosphere, but here the range of  $H_{LF}$  values obtained is still larger, with completely unrealistic temperatures that would be derived from that in the single case.

## 5 Conclusions

Two months of LF absolute reflection height data has been used to derive scale height estimates in the lower E region. These data have been qualitatively compared to meteor radar temperatures. The results are clearly qualitative. In addition, the influence of the extraordinary component of the LF wave on the results has to be analysed carefully and removed from the data set. This, however, makes the analysis somewhat complicated, and increases the number of potential errors.

Nevertheless, the long-period variations of both parameters available after low pass filtering are in adequate agreement, indicating that the temperature effect on the diurnal scale height change is inferable using the 177 kHz measurements at Collm, with short propagation distances. Since LF height measurements are available for more than 2 decades, this provides the opportunity to analyse planetary waves and their long-term variations and trends.

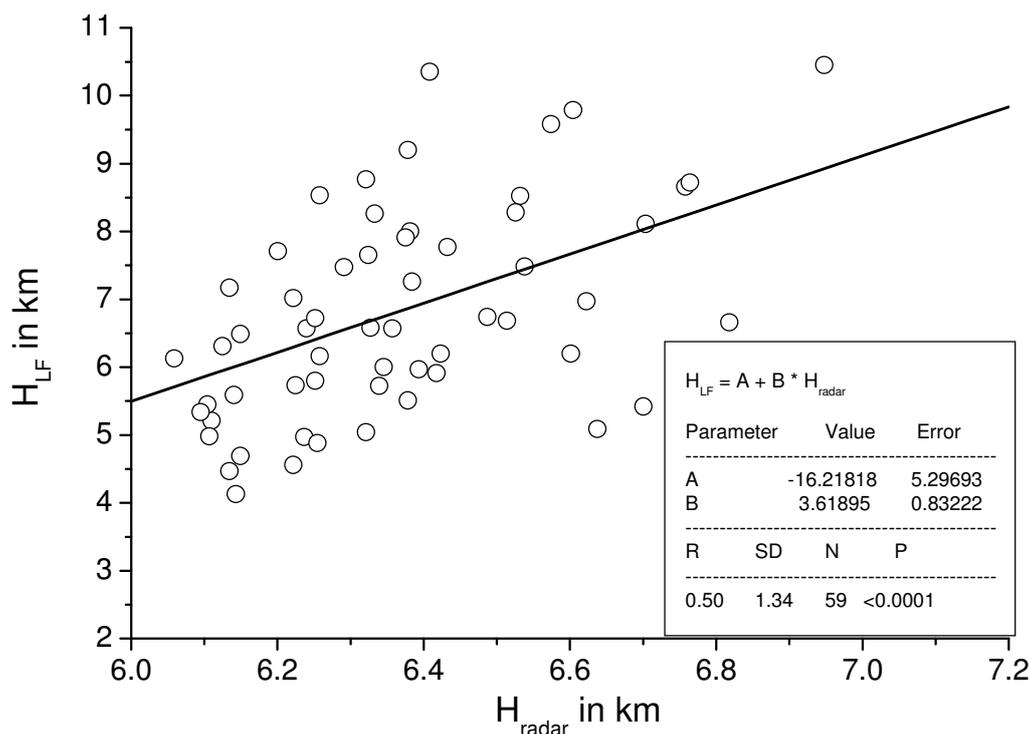


Figure 6. Scale height estimates  $H_{LF}$  calculated from LF reflection height changes vs. scale heights  $H_{radar}$  calculated from meteor radar.

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